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Airborne measurements of nucleation mode particles I: coastal nucleation and growth rates

C. D. O'Dowd¹, Y. J. Yoon^{1,2}, W. Junkerman³, P. Aalto⁴, M. Kulmala⁴, H. Lihavainen⁵, and Y. Viisanen⁵

¹Dept. of Physics & Environmental Change Institute, National Univ. of Ireland, Galway, University Road, Galway, Ireland

²Korea Polar Research Institute, KORDI, 7-50, Songdo-dong, Incheon 406-840, Korea

³Fraunhofer Inst. for Atmospheric Environmental Research, Kreuzteckbahnstr. 19, 82467 Garmisch-Partenkirchen, Germany

⁴Department of Physical Sciences, P.O.Box 64 (Gustaf Hållströmin katu 2) PL 64, 00014 University of Helsinki, Finland

⁵Finnish Meteorological Institute, Erik Palménin aukio, P.O.Box 503, 00101, Helsinki, Finland

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Abstract. A light aircraft was equipped with a bank of Condensation Particle Counters (CPCs) (50% cut from 3–5.4–9.6 nm) and a nano-Scanning Mobility Particle Sizer (nSMPS) and deployed along the west coast of Ireland, in the vicinity of Mace Head. The objective of the exercise was to provide high resolution micro-physical measurements of the coastal nucleation mode in order to map the spatial extent of new particle production regions and to evaluate the evolution, and associated growth rates of the coastal nucleation-mode aerosol plume. Results indicate that coastal new particle production is occurring over most areas along the land-sea interface with peak concentrations at the coastal plume-head in excess of 10^6 cm^{-3} . Pseudo-Lagrangian studies of the coastal plume evolution illustrated significant growth of new particles to sizes in excess of 8 nm approximately 10 km downwind of the source region. Close to the plume head ($<1 \text{ km}$) growth rates can be as high as $123\text{--}171 \text{ nm h}^{-1}$, decreasing gradually to $53\text{--}72 \text{ nm h}^{-1}$ at 3 km. Further along the plume, at distances up to 10 km, the growth rates are calculated to be $17\text{--}32 \text{ nm h}^{-1}$. Growth rates of this magnitude suggest that after a couple of hours, coastal nucleation mode particles can reach significant sizes where they can contribute to the regional aerosol loading.

Marine aerosol production and dispersion is of considerable importance globally given the large oceanic coverage of the Earth's surface and the underlying low albedo of the ocean surface. In particular, small changes in the availability of marine stratiform cloud nuclei availability can have profound impacts in cloud micro-physics and radiative properties (Slingo, 1999). Various components of the marine aerosol have been implicated in natural feedback processes involving the marine biota, aerosol production and aerosol interactions with clouds and radiation (Charlson et al., 1987; O'Dowd et al., 2004). Despite significant effort, new particle production (via homogeneous nucleation) over the ocean is still a poorly understood phenomenon. Considerable focus in recent years has centred on the production of new particles from the homogeneous nucleation of sulphuric acid, water vapour and ammonia with the sulphuric acid being derived from oxidation products of dimethylsulphide (DMS). Although there is some evidence for DMS-derived sulphuric acid being involved in the particle production process (Clarke et al., 1998), the most robust modelling studies have shown that while nucleation of sulphuric acid can occur, there is insufficient sulphuric acid available to grow the nucleated clusters to detectable sizes at larger than 3 nm (Pirjola et al., 2001; O'Dowd et al., 2003).

1 Introduction

The contribution of natural aerosol particle production to the global aerosol budget is important to quantify since aerosol radiative forcing provides the greatest uncertainty to current and future climate predictions (Houghton et. al., 2001).

Correspondence to: C. D. O'Dowd
(colin.odowd@emas.demon.co.uk)

The most intense and most frequent particle production events have been observed in coastal regions, presumably where biogenic emissions are enhanced. As far back as 1897, John Aitken (Aitken, 1897) reported elevated concentrations of small particles in coastal air along the Scottish coastline and attributed these to shoreline emissions. The phenomenon of coastal particles has been studied in detail with the aim of elucidating processes likely to be occurring over the open ocean since the strength and frequency of these events were

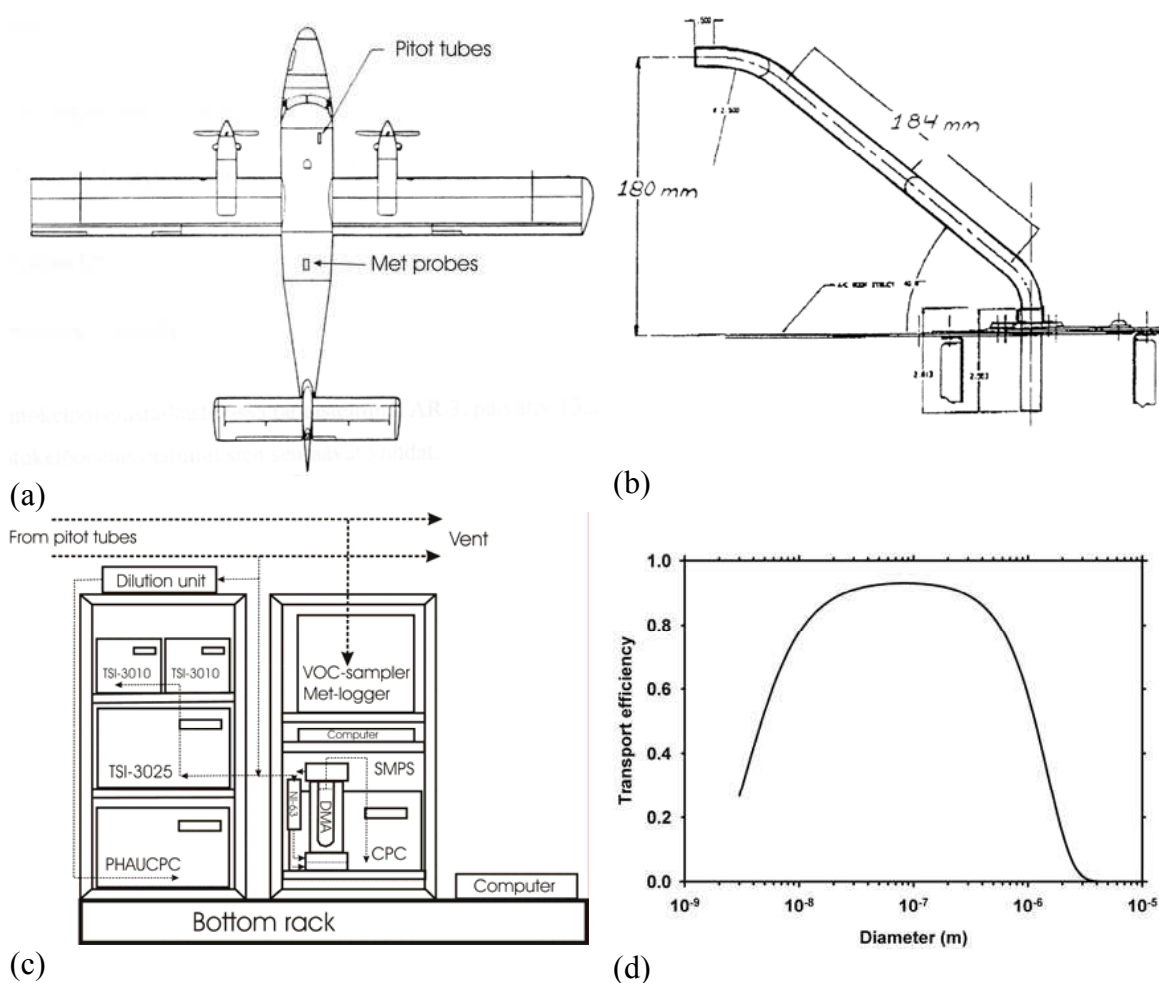


Fig. 1. Schematic of the GTK Twin otter and location of Pitot and met-probe tubes; **(b)** schematic of aerosol Pitot inlet tube; **(c)** schematic of aerosol instrumentation rack; **(d)** calculated aerosol losses of air speed of $0.75 \times 55 \text{ m s}^{-1}$.

high. O'Dowd et al. (1998) linked the occurrence of particle production events at Mace Head to enhanced biogenic emissions under low tide conditions and concluded that they were photochemically-driven (O'Dowd et al., 1999). It was postulated that biogenic iodine oxides were the most likely contributor to coastal nucleation events, although the exact mechanism still remains a topic for debate (McFiggans, 2005; O'Dowd et al., 2004).

The PARFORCE (Particle Formation and Fate in the Coastal Environment) project (O'Dowd et al., 2000) focused on the elucidation of the key processes driving coastal nucleation events and reported that peak particle concentrations could exceed 10^6 cm^{-3} and that formation events could last up to 8 hours under favourable conditions. The rate for formation of 3 nm particles was estimated to be of the order of $10^4 \text{ cm}^{-3} \text{ s}^{-1}$ with nucleation rates being of the order of $10^6 \text{ cm}^{-3} \text{ s}^{-1}$. Dal Maso et al. (2002), calculated growth rates of the order of 15–180 nm per hour for typical events

around Mace Head. Growth rates of this magnitude suggest that coastal particles could reach cloud condensation nuclei sizes (typically $>100 \text{ nm}$) in a matter of hours. The aforementioned growth rates were typically calculated about 100 m from the tidal plume head with minimal dilution of the plume having taken place. To provide a more accurate estimate for the growth rates of new particles in coastal aerosol plumes, airborne measurements need to be applied. O'Dowd (2002) conducted airborne measurements in the vicinity of Mace Head onboard the UK Met. Office C-130. During that study, particle production events were observed along the Irish – Atlantic coastline and the coastal plume was observed to extend about 250 km downwind. During these studies, both the scattering coefficient and the CCN concentration was observed to increase in the plume; however, no growth rates were calculated. In this study, we present improved estimates of coastal aerosol growth rates in the early stages of the development of coastal aerosol plumes on the

North East Atlantic coastline using a light aircraft and an aerosol payload designed to capture microphysical changes in the aerosol size distribution at sizes below 10 nm.

2 Experimental

The project formed the 1st of a series of 3 QUEST intensive field studies aimed at quantification of nucleation in the European boundary layer (Laaksonen et al., 2003). The project was based at Mace Head on the Irish Atlantic coastline. A complete summary of the station and facilities is found in Jennings et al. (2003). The station is subjected to regular clean marine air and many studies have shown that nucleation bursts occur regularly in this location (O'Dowd et al., 1998, 1999). The nucleation events have been linked to the increased emission of biogenic vapours released from macro algae exposed under low tide conditions (O'Dowd et al., 2002a, b).

2.1 Twin Otter

The measurement platform was a DHC-6/300 Twin Otter STOL fixed-wing, twin engine aircraft OH-KOG. The plane was manufactured in Canada by De Havilland Canada 1979. The aircraft is owned and maintained by Finnair and operated by a subsidiary of Finnair, Malmilento. The Geological Survey of Finland is responsible for all the research flights of the aircraft. It has been mainly used for aerogeophysical measurements but for these studies, the payload was modified for aerosol and meteorological measurements. A schematic of the aircraft and location of the aerosol inlets and meteorological probes are shown in Fig. 1. Typical air speed was of the order of 55 m s^{-1} .

2.2 Aerosol measurements

The aerosol sample for the instruments was taken through a pitot tube made of Stainless Steel installed to the roof of the airplane, just behind the pilot's seat on the right side of the plane. The pitot tube inner diameter was 10 mm and it extended 20 cm out from the roof (see Fig. 1). Inside the plane the pitot tube was connected to the Stainless Steel sample line 10 mm of diameter with a straight-angle fitting. The sample line made again a straight angle turn and the sample for the aerosol instruments was taken from the main line through a T-fitting with a Stainless Steel tube of 4 mm inner diameter. The flow of the main sample line was generated using the dynamic pressure of the pitot tube. The outlet of the main sample line was located under the plane. The total length of the sample line was 2.8 m to a distribution point, following which, a further 2.5 m was used to provide flow to the furthest most instrument from the distribution point. The flow velocity in the main sample line was of the order of $30\text{--}40 \text{ m s}^{-1}$.

The aerosol setup consisted of three condensation particle counters, one TSI model 3025 with cut off diameter

around three nanometers and two TSI model 3010 counters with measured 50% lower cut-off diameters of 5.4 and 9.6 nanometers, a SMPS (Scanning Mobility Particle Sizer) measuring particle size distribution between three and fifteen nanometers. A dilution system was applied for CPC measurements, with a ratio of 1 to 7.

The flow speed in the aerosol sampling duct was approximately 0.75 of the true air speed. Particle losses are calculated for the envelop of sample velocities corresponding to the aircrafts air speed and 50% of the air speed. Included in the calculations were losses due to an-isokinetic sampling, diffusion losses, gravitational losses and impacts in duct bends. It was found that the air speed did not make significant differences to the calculated losses and that the flow was turbulent for the flow regimes used. Particle losses were minimal for sizes of 20 nm to 150 nm and increased to 65% at sizes of 3 nm compared to 30% at 10 nm (See Fig. 1).

The concentration of the three CPCs were checked when no nucleation mode was presented and absolute concentrations were normalized to the 10 nm CPC. With this system we were able to record the total particle concentration using the TSI model 3025 data and by subtracting the readings between the counters, we got an estimate of the ultrafine particle concentration between 3 and 5.4 nm and 5.4 and 9.6 nm (i.e. the 50% cut-size for the three CPCs used). Since the detection efficiency curves are not descending sharply around the cutoff diameters, the calculated ultrafine particle concentrations are just estimates for the concentration in each size class. The three particle counter readings were recorded once a second.

The SMPS system consisted of TSI model 3025 condensation particle counter, Hauke-type DMA (Differential Mobility Analyzer) and a 370 MBq Ni-63 aerosol neutralizer. The sheath flow rate of the DMA was 15 lpm and aerosol flow rate 1.7 lpm. The sheath flow rate was controlled with a critical orifice and arranged as closed loop. The sheath flow rate was dried. A portable computer and a National Instrument NI DAQCard-6036E was used to scan the DMA voltage, record the voltage, time and particle count from the CPC with ten hertz frequency. Also temperature, humidity and pressure were recorded once during the scan. The scan time was typically thirty seconds. The data was inverted online and both raw data and inverted data were saved to the disk. In the SMPS inversion the measured delay time (1.25 s) from the DMA to the CPC was taken into account but the time constant due to the sample "smearing" in the DMA, sampling lines and the CPC was not taken into account because the time constant was fairly small, some tenths of seconds.

2.3 Gas measurements

The gas sample was taken through a similar pitot tube that aerosol sample. The inlet was located behind the aerosol sample inlet and slightly off-shifted from the aerosol. Inside the plane, the pitot tube was connected to the Teflon

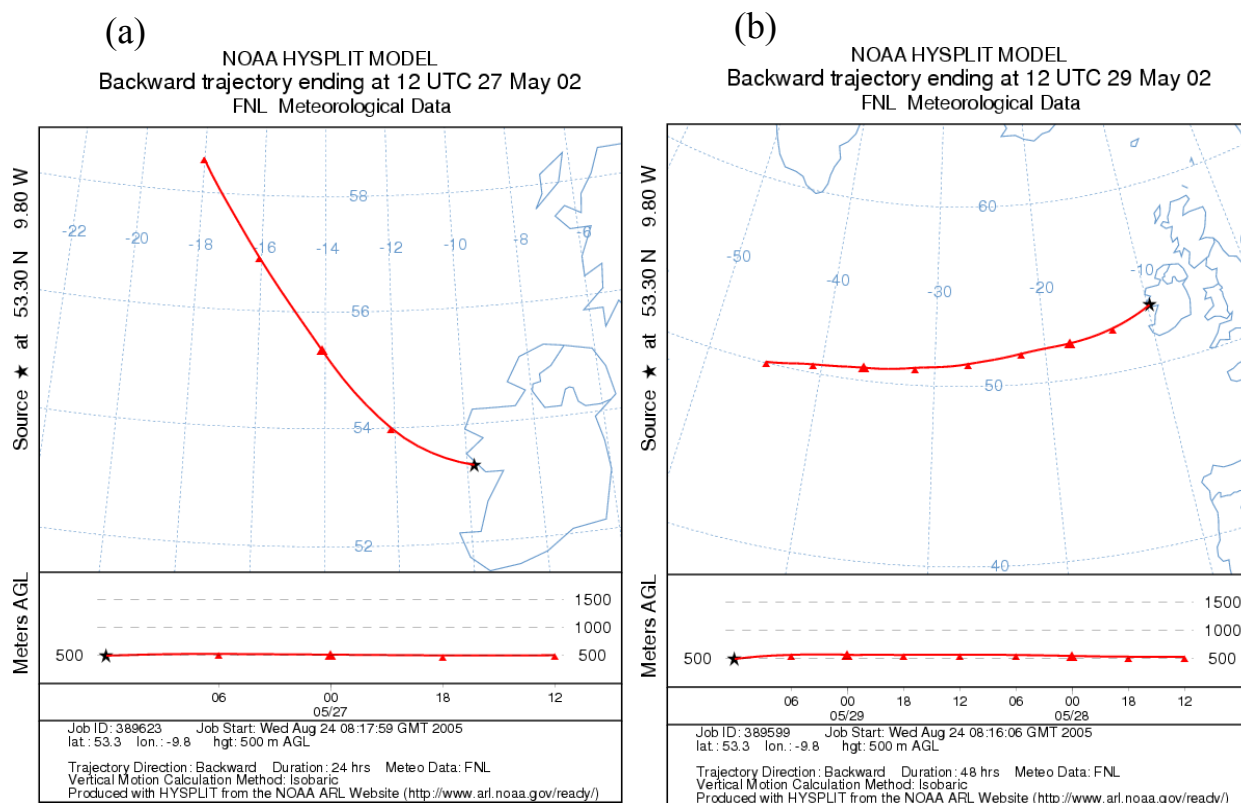


Fig. 2. Air mass back trajectories arriving at Mace Head for (a) on 27 May, 2002 and (b) on 29 May, 2002.

sample line 10 mm of diameter with a straight-angle fitting. The total length of the sample line prior gas samplers was 3 m. The flow of the main sample line was generated using the dynamic pressure of the pitot tube. The outlet of the main sample line was located under the plane. The flow rate in the main sample line was unknown.

A special holder unit was constructed for Tenax tube sampling. The Teflon sample line was connected to 12 mm outer diameter and about 35 cm long Stainless Steel tube. Ten 6 mm outer diameter and 2 cm long tubes were welded into this tube. Tenax sampling tubes were connected into these with stainless steel Swagelok fittings. Ozone remover grids were located in the fittings prior the Tenax tubes. Each tube had a valve after it. The valves were connected to same vacuum line. The flow through the sample tube was controlled with a critical orifice. The flow rate through the Tenax tubes was 260 cm³/min. The holder unit was removable so that Tenax tubes could be changed in clean environment. Typical sampling time was about 10 min. From Tenax tubes monoterpenes and aromatic hydrocarbons were analysed with gas chromatograph with mass selective detector.

For flask samples a pump was connected into the main gas sample line with a Teflon tubing. Pump provided about 1 bar over pressure into to flasks. The time to fill a flask

was about 30 s. Light hydrocarbons (C₂–C₆) were analyzed from flask samples. Unfortunately, the gaseous species were below detection on the Mace Head flights and results of these samples over the boreal forests are reported in Part II of these paper.

2.4 Meteorological measurements

Sensors for temperature, relative humidity and solar radiation intensity were located on the roof of the airplane, about in the middle of the whole plane. The pressure sensor was located inside the airplane. Temperature and relative humidity was measured with Vaisala HMP45 D probe. The probe was shielded against direct solar radiation. Pressure was measured with Vaisala PTB101B analog barometer. Total solar radiation was measured with Licor, model LI-200 SZ, pyranometer sensor. Position of the airplane was recorded with Garmin GPS receiver with antenna on the roof of the airplane. Temperature, relative humidity, solar radiation, pressure and position were recorded every second.

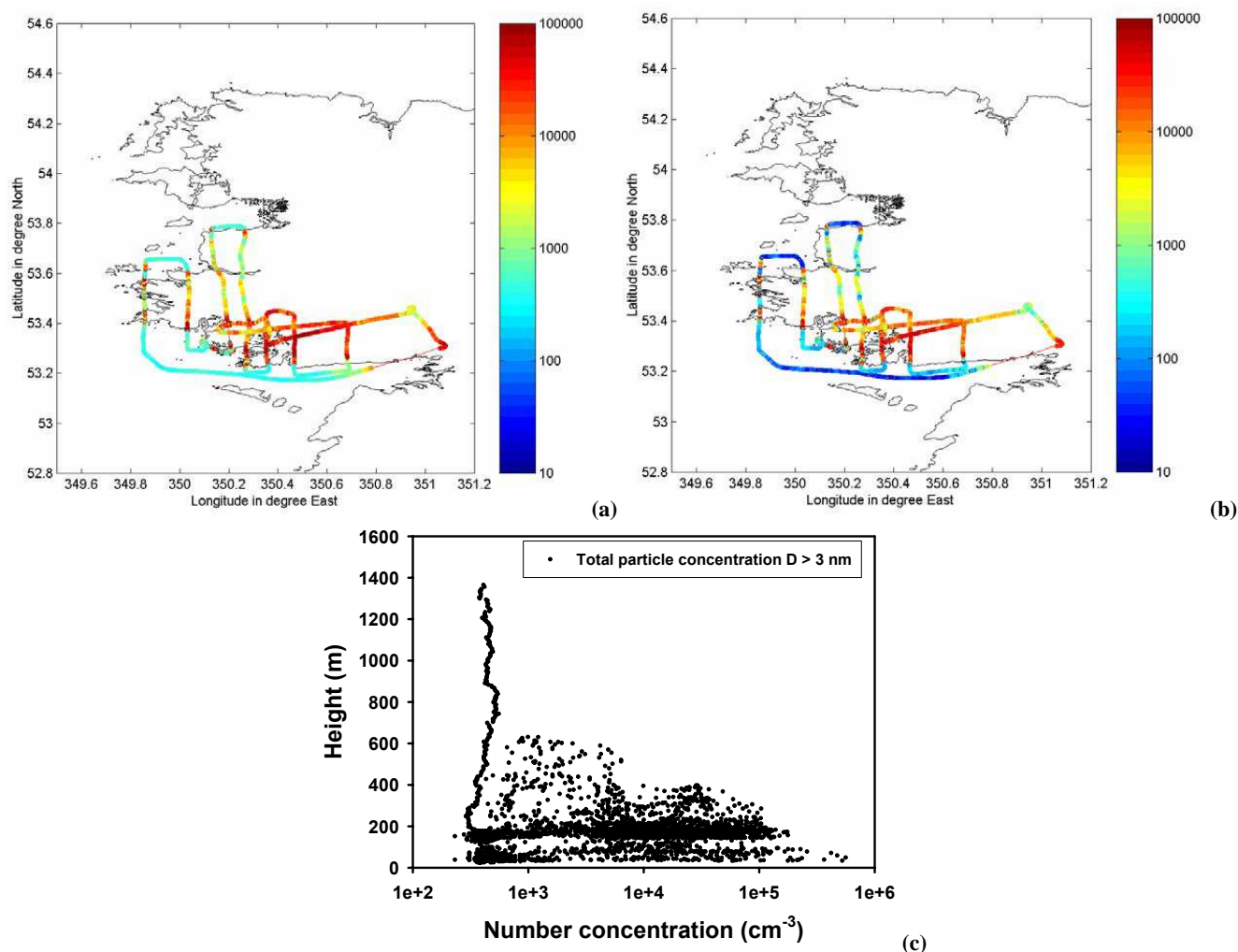


Fig. 3. (a) Total particle concentration for $d > 3$ nm during flight on 27 May 2002. (b) same as (a) except for total particle concentration at sizes $3 < d < 10$ nm, and (c) Scatter plot of total particle number concentration ($D > 3$ nm) versus height from the mapping flight on 27 May 2002. Concentrations units are particles cm^{-3} .

3 Results

A combination of mapping flights and Lagrangian flights were conducted. The mapping flights focused on locating coastal hot-spots for aerosol production along the coastal zone while the Lagrangian flights were aimed at quantifying aerosol growth in plumes emerging from the hotspots. Two flights are presented in this study.

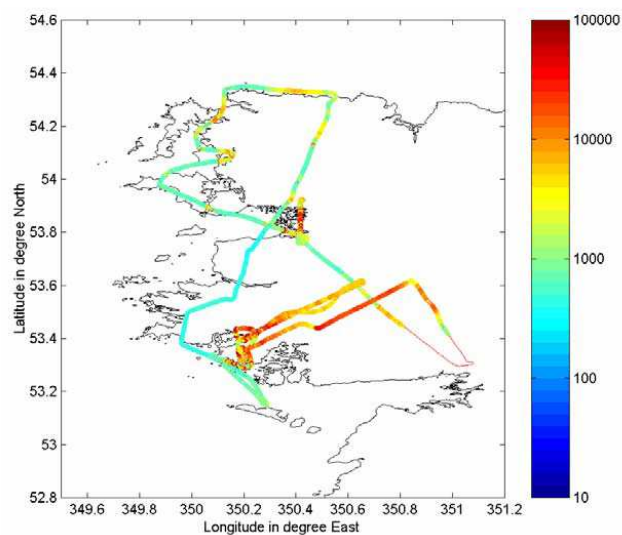
3.1 Mapping flight

The mapping flight was conducted around the coastal regions of Mace Head on the 27 May 2002. Air mass back trajectories, calculated from HYSPLIT (Draxler et al., 2003), for this day indicate a clean westerly air flow arriving at the station (Fig. 2a).

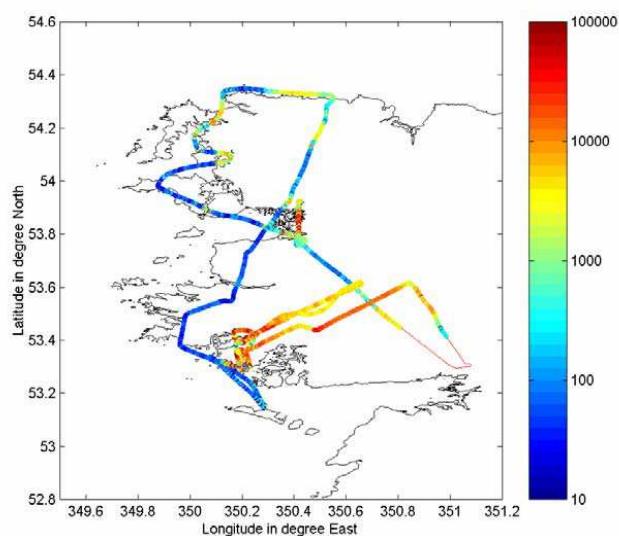
The aircraft flight departed from Galway, and made way to the Mace Head region over the ocean and clear of any coastal aerosol sources (Fig. 3). The total aerosol concentration was of the order of 500 cm^{-3} in the regions of air flow not advecting over tidal areas. During this leg, the maximum height the aircraft reached was 1500 m above sea-level (and in the free troposphere) where the total particle concentration was less than 400 cm^{-3} (Refer to Fig. 3c showing scatter plot of total particle concentration for various heights). After reaching the western-most region of the coastline, the aircraft proceeded to map aerosol concentration distributions at 50 m above sea/ground-level. The mapping legs were conducted north-south and relocating east and downwind after each transect. Total particle concentration and the concentration of particles between 3 nm and 10 nm are shown in Fig. 3. A clear increase in particle concentration is seen

Table 1. Short range Lagrangian run scan information (time, geographical location, log-normal fit modal diameter standard deviation, and modal concentration).

	Time JD	Location	Mean Diameter (nm)	Sigma	Total
Production area		053.390 N, 009.798 W	1.0	–	–
SD1	149.6189	053.395 N 009.789 W	4.49	1.122	41220
SD2	149.6187	053.401 N 009.780 W	5.44	1.123	12342
SD3	149.6184	053.409 N 009.766 W	6.52	1.128	8063



(a)



(b)

Fig. 4. (a) Total particle concentration for $d > 3$ nm during flight on 29 May 2002, and (b) same as (a) except for total particle concentration at sizes $3 < d < 10$ nm. Concentrations units are particles cm^{-3} .

in all regions downwind of the coastline. Peak concentrations reached $473\,000\text{ cm}^{-3}$ (not accounting for particle losses averaging approximately 50% at sizes less than 10 nm) with the strongest signal seen in the region to the south west (Mweenish area) and North (Roundstone area) of Mace Head. The concentrations of particles larger than 10 nm measured from CPC 3010 over these regions remained less than $20\,000\text{ cm}^{-3}$, indicating almost all particles detected are recently formed. These results indicate that there are many sources of new particles along the coastline in the vicinity of Mace Head. In fact, significantly stronger regions other than Mace Head were identified, the strongest being Mweenish and Roundstone. A similar pattern was seen on other flights (see next section) and previously by O'Dowd (2002).

3.2 Lagrangian flights

During the 29 May, the air mass trajectories indicate a south westerly air flow, again in clean marine air (Fig. 2b). As stated above, the strongest source regions were identified as Roundstone and Mweenish. The Roundstone area, to the north-north west of Mace Head, was selected as the starting point of a Lagrangian flight to quantify growth rates of new particles in the coastal plume. The Lagrangian flight was actually a combined mapping and Lagrangian flight with mapping performed over a more extensive area compared to the first mapping flight.

At the Mace Head station, there was only very weak evidence of a nucleation burst with peak particle concentrations only briefly exceeding $10\,000\text{ cm}^{-3}$. A similar pattern was seen in the mapping component of the flight with only a few regions exhibiting elevated particle concentrations (Fig. 4). The intensity of nucleation events during this week was considered considerably low by comparison to similar events at the same time of year (O'Dowd et al., 2002a) where particle concentrations were of the order of 10^6 cm^{-3} .

A strong formation of particles was detected over Roundstone area (53.39° N , 9.80° W) where seaweeds are widely found. Meteorological wind conditions for this day were 220° direction and speed of 8 ms^{-1} , respectively. The aircraft flew towards the particle production location (*Lag-run 1*) maintaining 220° flying direction, then turn around and

Table 2. Calculated aerosol growth rates from the short range Lagrangian run.

Scan number	Distance from production area (m)	Air mass traveling time from production area (h)	Growth rate from 1.0 nm to modal diameter (nm h^{-1})	Growth rate of particles of which diameter is larger than the modal diameter by 30% (nm h^{-1})
SD1	816.5	0.028351	123.1	170.6
SD2	1710.8	0.059403	74.7	102.2
SD3	2997.4	0.104076	53.0	71.8

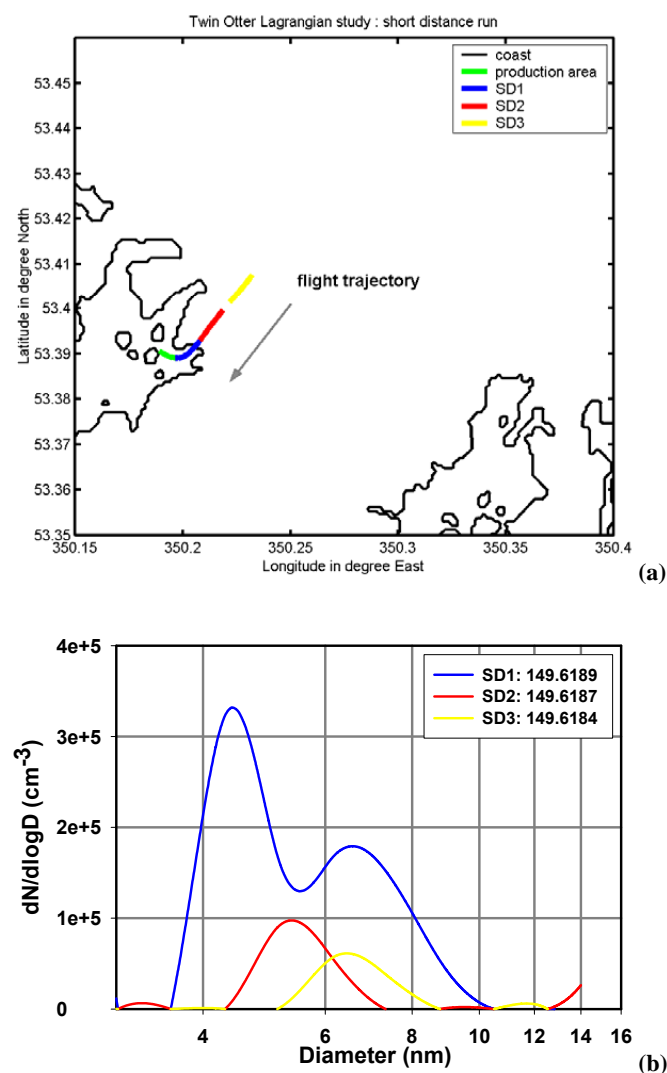


Fig. 5. (a) Trajectory of Lag-run 1 to the particle formation region (grey arrow). The selected particle formation regions are represented in green lines. Blue (SD1), red (SD2), and yellow (SD3) lines correspond to the distance covered by each 20 seconds SMPS scan, with SD1 closest to the source region, and (b) SMPS size distributions for scans over regions SD1, SD2 and SD3. For SD1, only the first modal diameter was used for growth rate analysis.

flew along the wind direction from the source location (*Lag-run 2*).

3.2.1 Lag-run 1: short distance Lagrangian run

For *Lag-run 1*, three spectra were selected to extract growth rate of recently formed particles assuming that particles were formed at production location which is colored as green in Fig. 5. In this figure different colors represent the flying trajectory of the aircraft covering range of 3 individual SMPS scans and aerosol number size distributions are presented in Fig. 5. The scanning starting time and location, and log-normal fitting values for each scan are summarized in Table 1. This table also includes location of source region. The mean nucleated particle, or cluster, diameter in the production zone was assumed to be 1.0 nm.

To calculate growth rates of the coastal nucleation mode, the distances from the production location to each of the scan area mid-points were calculated, then the time needed for the air mass to travel from the source region to each scan locations was derived by assuming a wind speed of 8 ms^{-1} . Growth rates of the recently formed particles were calculated by assuming an initial diameter of 1.0 nm and then quantifying the modal diameter for the appropriate scans. The distance from the production area, time travelled from the production area and the growth rates are presented in Table 2. Two growth rates are presented – one for the modal diameter growth and another for the growth rate of particles with diameter 30% larger than the modal diameter. A bimodal size distribution was observed from the SD1 scan. The second size mode around 7 nm was not considered in the growth rate calculation, which is believed to grow into the larger sizes than the SMPS scanning range.

Closest to the source (817 m), a growth rate of 123 nm h^{-1} was observed, reducing to 75 nm h^{-1} at 1710 m and to 53 nm h^{-1} at 2997 m. From these results, it can be observed that significant particle growth is observed close to the plume head, in agreement with previous ground based studies, and that this growth rates reduce with distance from the source indicating that gases responsible for condensation growth are diluted.

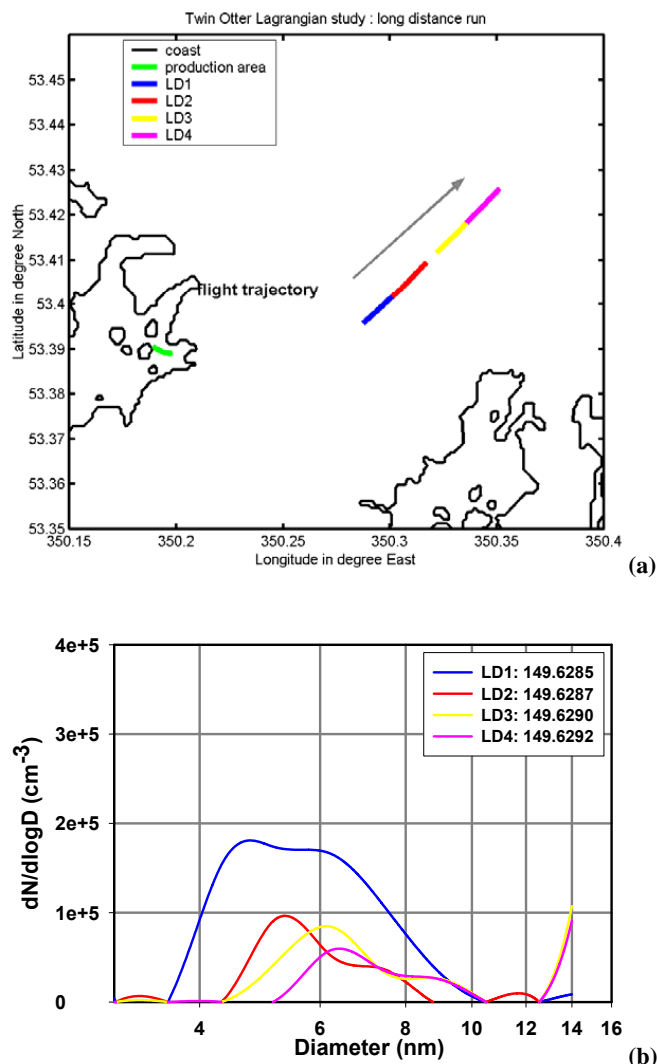


Fig. 6. (a) Trajectory of Lag-run 2 from the particle formation region (grey arrow). The selected particle formation regions are represented in green lines. Blue (LD1), red (LD2), yellow (LD3) and purple (LD4) lines correspond to the distance covered by each 20 seconds SMPS scan, with LD1 closest to the source region, (b) SMPS size distributions from scans corresponding to regions LD 1–4. Each scan has 2 modal diameters, so growth rates were separately calculated for 1st and 2nd modal diameters.

3.2.2 Lag-run 2: long distance Lagrangian run

From Lag-run 2, four spectra were selected. The production location was assumed to be same as Lag-run 1. As is seen in Fig. 6, four successive scans were done relatively far from the source location and size spectra possess 2 modes and 2 modal diameters, suggesting perhaps more than one source region along the trajectory or nearby plumes of differing condensable vapour source strengths mixing. For the analysis presented here, the latter assumption is made. Tables 3 and 4 summarize physical properties of each scan

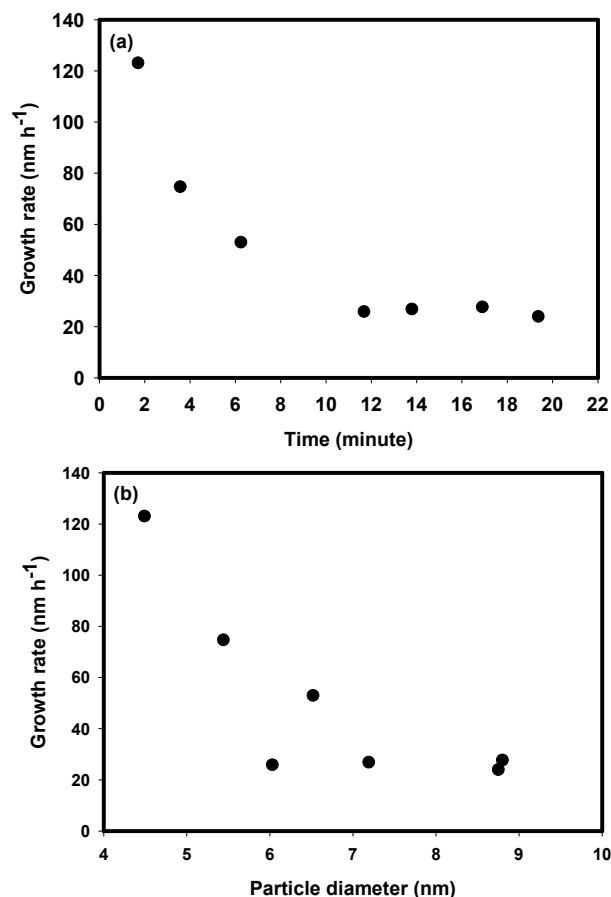


Fig. 7. Growth rates as a function of (a) air mass aging time and (b) particle modal diameter.

and growth rate for 2 modal diameters, in a same manner as Lag-run 1. At a distance of 5600 m, the growth rate of mode 1 is 19.4 nm h^{-1} while the growth rate for mode 2 is 25.9 nm h^{-1} . For the growth rates of diameters 30% larger than the modal diameter, the respective values are 26.8 and 35.1 nm h^{-1} . The growth rates for both modes again gradually decrease with decreasing distance and the rate for mode 1 at a distance of 9298 m was 17.0 nm h^{-1} , while for mode 2, it was 24.0 nm h^{-1} . These results illustrate that even at a distance of 10 km from the source, the nucleation mode growth rate is significant, despite dilution. Although a gradual decline of growth rate occurs with increasing distance, at these distances the reduction in the growth rates is becoming minimal.

3.2.3 Combination of short and long distance Lagrangian run

Growth rates derived during the two types of Lagrangian run showed a decreasing pattern as the air mass age. The change of growth rate as a function of time and modal diameter in an

Table 3. Long range Lagrangian run scan information (time, geographical location, log-normal fit modal diameter standard deviation, and modal concentration).

	Time JD	Location	1st mode		2nd mode			
			Mean Diameter (nm)	Sigma	Total	Mean Diameter (nm)	Sigma	Total
Production area		053.390 N 009.798 W	0.5	–	–	–	–	–
LD1	149.6285	053.395 N 009.714 W	4.78	1.156	29766	6.03	1.248	40511
LD2	149.6287	053.401 N 009.700 W	5.38	1.130	12706	7.19	1.100	4143
LD3	149.6290	053.411 N 009.681 W	6.10	1.154	13198	8.80	1.092	2646
LD4	149.6292	053.418 N 009.666 W	6.50	1.128	7815	8.75	1.092	2620

Table 4. Calculated aerosol growth rates from the long range Lagrangian run.

	Distance from production area (m)	Air mass traveling time from production area (h)	1st mode		2nd mode	
			Growth rate from 1.0 nm to modal diameter (nm h^{-1})	Growth rate of particles of which diameter is larger than the modal diameter by 30% (nm h^{-1})	Growth rate from 1.0 nm to modal diameter (nm h^{-1})	Growth rate of particles of which diameter is larger than the modal diameter by 30% (nm h^{-1})
LD1	5603.8	0.194576	19.4	26.8	25.9	35.1
LD2	6619.2	0.229833	19.1	26.1	26.9	36.3
LD3	8109.4	0.281576	18.1	24.6	27.7	37.1
LD4	9298.0	0.322847	17.0	23.1	24.0	32.1

air mass are shown in Fig. 7. The increasing 2nd modal diameters obtained during the long-range Lagrangian run were not taken into account because the source region of these modes cannot be defined as clearly as the first modes. It was found that growth rate of freshly formed particles decrease dramatically as the air mass ages. The growth rate of freshly formed particles was higher than 100 nm h^{-1} , and this rate decrease rapidly until the saturation point achieved in 12 min showing the rage of 20 nm h^{-1} – 40 nm h^{-1} which is equivalent to 7 nm diameter.

The decrease of the growth rate as a function of size seems to contradict with the recent observation by Kulmala et al. (2004b) showing that growth rate will increase as a function of size. However, this is not the case, since here the dilution will decrease both particle concentration (decreasing growth due to reduced coagulation) and vapor concentration (decreasing the growth due to reduced condensation).

There have been a number of works on the growth rates of freshly formed particles, and these results were summarized in Kulmala et al. (2004). Kulmala et al. (2004) surveyed published data on the particle formation and reported that typical particle growth rates range from 1 nm h^{-1} to 20 nm h^{-1} depending on the place of particle formation and condensing vapour availabilities. Particle growth rates at rural and urban areas are typically smaller than 10 nm h^{-1} (Shi et al.,

2001; Weber et al., 1997; Birmili et al., 2001; Kulmala et al., 1998; Aalto et al., 2001). Growth rates at polar regions usually range from 0.1 to 2 nm h^{-1} even under a polar coastal environment due to low temperature and lack of condensing vapour gases (Gras, 1993; Park et al., 2003; Ito, 1993). Previous study by O'Dowd et al. (2002) and Dal Maso et al. (2002) reported growth rate of recently formed particles as 15 – 180 nm h^{-1} at Mace Head based on the in-situ measurements during PARFORCE campaign. It is mostly likely that particles formed at coastal environment can grow much faster than any other other locations due to the abundance of condensing vapours, and pseudo-Lagrangian study of our work also supports this hypothesis. Our Lagrangian study showed that freshly formed particles in a same air mass can grow rapidly when 3–4 nm particles occupy majority of number concentrations. This rapid growth is hindered by the consumption of condensing or nucleating vapours, due to increasing particle surface area. Nevertheless, the growth rates for 7–8 nm diameter particles are still around 20 – 40 nm h^{-1} , showing significantly larger value than those observed in continental particle production event.

4 Conclusions

Two flights mapping particle formation along the Irish Atlantic coastline and studying aerosol evolution in the coastal plume are presented. It was found that formation is almost ubiquitous along the Irish coastline with particle concentrations exceeding 10^6 cm^{-3} in the coastal plume head where the mean particle size is about 3 nm. The formation is observed inside the boundary layer. As the coastal plume advects downwind (approximately 50 km), the mean size increases from 3 nm to >10 nm. Close to the plume head (<1 km) growth rates can be as high as $123\text{--}171 \text{ nm h}^{-1}$, decreasing gradually to $53\text{--}72 \text{ nm h}^{-1}$ at 3 km. Further along the plume, at distances up to 10 km, the growth rates are calculated to be $17\text{--}32 \text{ nm h}^{-1}$. Growth rates of this magnitude suggest that after a couple of hours, coastal nucleation mode particles can reach significant sizes where they can contribute to the regional aerosol loading.

The observed growth rates in the coastal environment are significantly higher than in any other environments. However, after couple of minutes (kilometers) the growth seems to reduce at the level of 20 nm h^{-1} , which has been observed also in other (mainly urban) conditions. This high growth rate indicates high concentration of condensable vapors typically order of tens of ppt, and the reduction indicated dilution of condensable gases.

In future studies the coastal nucleation and subsequent growth should be parameterized to use this knowledge in air quality and global models. The further analysis would also include the analysis of condensation and coagulation sinks and the contribution of coagulation on growth processes.

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